

Frequency Modulation of 0S2-B

Herbert Weidner ^a

Abstract: Precision measurements of the 0S2 quintet after the 2004-12-26 earthquake show that the spectral line near 304.6 μHz is frequency modulated with a remarkably large frequency deviation of 0.18 μHz . By proper choice of the integration length, the center frequency can be determined with high precision.

Introduction

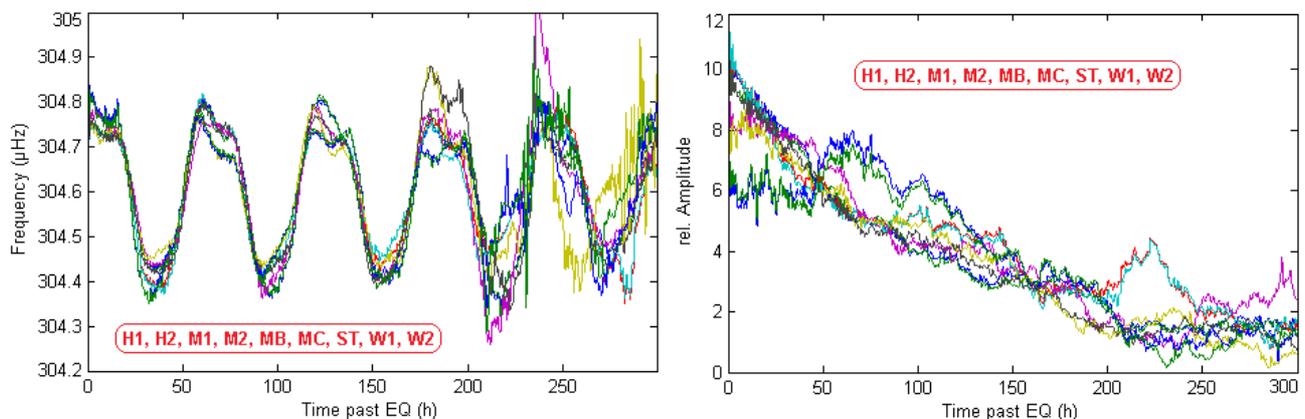
After earthquakes, the Earth vibrates like a bell at different frequencies. The lowest ones near 300 μHz are particularly interesting because of their relative proximity to the rotation frequency of the earth. The remarkably wide error bars of all previous measurements are probably caused by the overlooked frequency modulation of these natural frequencies. High precision can only be achieved when the integration period is adapted to the modulation frequency.

The Preparation of the data was described here^[1]. The essential points: Before calculating the frequency, each data segment passed a narrow band [Sinc filter](#) with the bandwidth 0.8 μHz . A broadening of the bandwidth hardly changes the results. However, a reduction causes severe distortion.

To distort the data as little as possible and to increase the frequency resolution, a new method^[2] has been developed, eliminating the need for a window function and zero padding. In this way, any data corruption by the arbitrary window function is avoided. FFT is replaced by the faster Goertzel algorithm, because this allows period lengths that are not a power of 2.

The Frequency of 0S2-B

The instantaneous frequency differs almost always from the average frequency. The striking [frequency deviation](#) ($\pm 0.18 \mu\text{Hz}$) far exceeds the deviation^[3] of the other lines of the $0S_2$ quintet^[4]. This frequency modulation was discovered (see Figure 5) a long time ago in a completely different way^[5]. At that time, the same lowest modulation period (60 hours) was measured. All European stations consistently show the same frequency modulation of the spectral line $0S_2$ -B. Due to the small mutual distances, the phase shift is nearly negligible.



Since the average frequency of a FM-oscillation is strongly influenced by the length of the integration period^[1], a random choice of the integration time will surely generate an incorrect result. The following table shows how the mean value depends on the integration time (first 210 hours after the earthquake). As shown below, the modulation period length is $T = 59.87$ hours.

a) 16. June 2015, email: herbertweidner@gmx.de

Int. Time	$1 \cdot T$	$1.5 \cdot T$	$2 \cdot T$	$2.5 \cdot T$	$3 \cdot T$	$3.5 \cdot T$	$4 \cdot T$
f in μHz	304.60231	304.62266	304.60057	304.61394	304.60008	304.61013	304.60112

If we restrict ourselves to the fields *without* systematic error (highlighted in yellow), the jackknife method provides the mean frequency (304.60102 ± 0.00048) μHz .

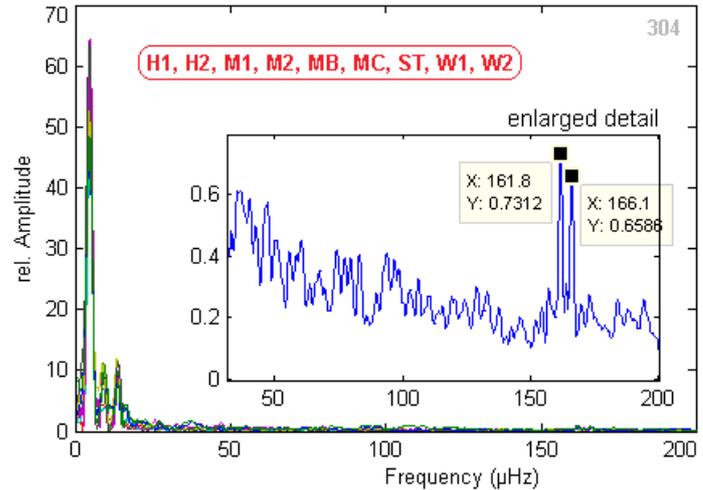
The Modulation Frequencies

Another way to calculate the precise average frequency is to compensate for the frequency deviation by an opposite phase signal. As with the "detiding", the amplitudes and phases of all significant modulation frequencies must be determined. The spectral line ${}^0\text{S}_2$ -B is modulated with two different frequency groups. The low group includes three individual frequencies: $f_1 = 4.64 \mu\text{Hz}$, $f_2 = 9.22 \mu\text{Hz}$ and $f_3 = 13.80 \mu\text{Hz}$. The higher group near $165 \mu\text{Hz}$ is much weaker and may be neglected.

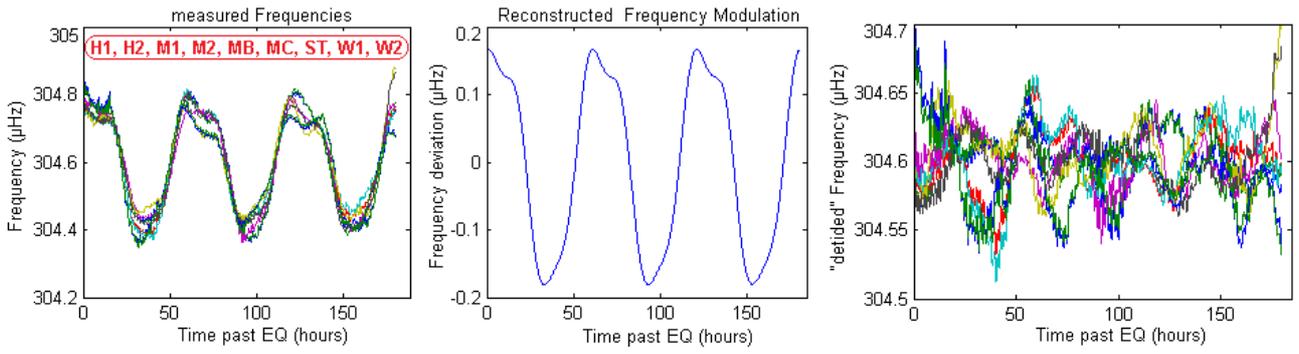
The strongest modulation frequency f_1 determines the period ($T = 59.87$ hours) of the very large frequency deviation.

With sine waves of these three frequencies, the actual time-dependent frequency course of each station can be reconstructed with high accuracy. The required amplitudes (in μHz) and phases ($0 \dots 2\pi$) are tabulated below.

Station	Ampl ₁	Ampl ₂	Ampl ₃	Phase ₁	Phase ₂	Phase ₃
F (μHz)	4.64	9.22	13.80	4.64	9.22	13.80
H1	0.18600	0.01710	0.02680	0.84	2.73	1.88
H2	0.19150	0.01780	0.02780	0.85	2.72	1.84
M1	0.17350	0.01050	0.02220	0.96	3.38	2.29
M2	0.16840	0.01090	0.02190	0.97	3.30	2.25
MB	0.18310	0.00760	0.02000	0.89	3.03	1.79
MC	0.16330	0.01920	0.03010	1.04	2.27	2.09
ST	0.19460	0.00850	0.03240	0.91	3.03	2.13
W1	0.16960	0.02170	0.02060	1.00	4.11	2.36
W2	0.17220	0.02240	0.02190	1.00	4.07	2.32

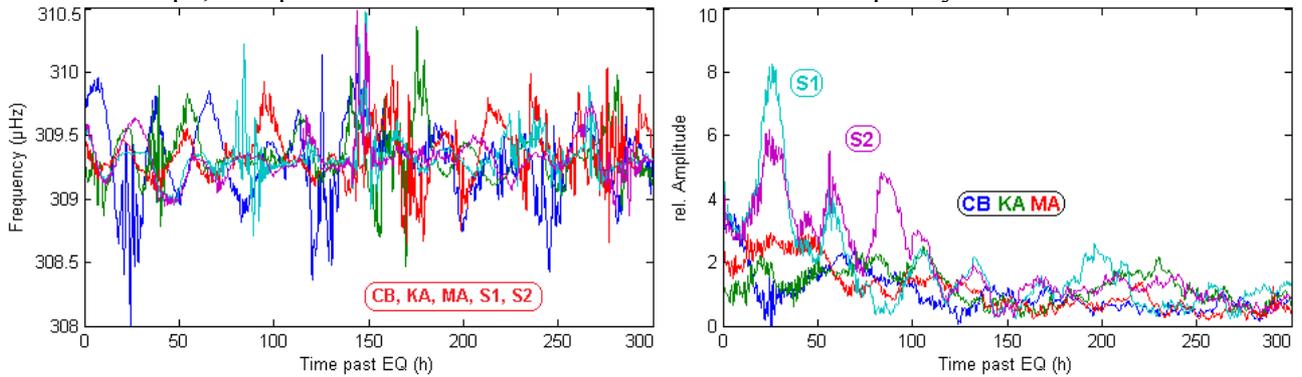


The modulation function of each European station is individually calculated (one example is shown in the middle picture below) and subtracted from the measured values, resulting in the right picture. Since only the three lowest and most intense modulation frequencies were compensated, the difference (shown in the right picture) is not just noise, but still shows some regular structures.



The average of the nine difference curves was calculated with the jackknife method, yielding $(304.60026 \pm 0.00027) \mu\text{Hz}$. In earlier measurements^[6], significantly larger error bands were given. This may have been caused by the ignorance of the frequency modulation with its consequences.

Outside Europe, the spectral line 0S_2 -B is too weak to allow for frequency determination.



Amplitude Decay and Q-Factor of 0S_2 -B

The amplitude reduction of the 0S_2 -B frequency $304.6 \mu\text{Hz}$ is expected to follow an exponential law. With the exception of W1 and W2, all the other curves of European SG stations match. Ignoring those two results, the decay during the first 220 hours past the earthquake may be described by the exponential function

$$A = A_0 \cdot e^{-\frac{t}{T}}$$

The time constant for the European stations (without W1, W2) is $T_{0S_2-B} = (138.25 \pm 4.97) \text{ hours}$.

The quality factor Q may be computed using the equation

$$A = A_0 \cdot e^{-\frac{t}{T}} \sin(\omega t + \varphi) = A_0 \cdot e^{-\frac{\omega t}{2Q}} \sin(\omega t + \varphi)$$

For $f_{0S_2-B} = 304.6 \mu\text{Hz}$, this equation yields

$$Q_{0S_2-B} = 476.3 \pm 17.1$$

Acknowledgments

Thanks to the operators of the GGP stations for the excellent gravity data. The underlying data of this examination were measured by a net of about twenty SG distributed over all continents, the data are collected in the Global Geodynamic Project^[7].

- [1] H. Weidner, Frequency Modulation of OS2-E, 2015, <http://viXra.org/abs/1506.0084>
- [2] H. Weidner, A New method for High-resolution Frequency Measurements, 2015, <http://vixra.org/abs/1506.0005>
- [3] The term *frequency deviation* is sometimes mistakenly used as synonymous with frequency drift, which is an unintended offset of an oscillator from its nominal frequency.
- [4] H. Weidner, Frequency Modulation of OS2-C, 2015, <http://viXra.org/abs/1506.0099>
- [5] H. Weidner, Envelope reconstruction of the OS2 – free oscillation of the earth, 2010, <http://arxiv.org/ftp/arxiv/papers/1012/1012.3259.pdf>
- [6] R. Häfner, R. Widmer-Schmidrig, Signature of 3-D density structure in spectra of the spheroidal free oscillation OS2, *Geophys. J. Int.*, 2012
- [7] The "Global Geodynamics Project", <http://www.eas.slu.edu/GGP/ggphome.html>